3D Inversion of Automated Resistivity Profiling (ARP) Data

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Electrical resistivity is a well established and widely used prospection method. The technique has also become a valuable and inseparable tool in archaeological site evaluation and excavation planning (Clark, 1990). Nowadays, it is routinely used in mapping the sub-surface resistivity variation of an archaeological site. The compilation of maps that transform the geophysical results into images that resemble a plane view of buried relics comprises the ultimate goal of geoelectrical prospection in archaeological areas (Scollar et al., 1986).

The standard mapping technique is sufficient and successful in most routine field applications (Tsokas et al., 1994). On the other hand, the use of a fixed-array length throughout the field survey hinders the interpretation of collected resistivity data in areas with more complex surface geology, as the current cannot penetrate to deeper levels in order to gain more depth information. The development of multiplexing resistance systems provided a step-forward in the determination of three dimensional distribution of apparent resistivity (Walker, 2000).

Additionally, the need for conducting high-resolution surveys characterized by a small measurement mesh and reducing as much as possible the survey cost without limiting the extent of the investigated area, has led to the development of mobile quadripoles allowing continuous measurement while moving (Panissod et al., 1997, 1998). These multi-electrode devices allow measurements over several depths of investigation in order to perform an exhaustive 3D exploration.

The Automatic Resistivity Profiling (ARP) method is a V-shaped multipole system with one transmitting dipole and three receiving dipoles, the length of which increases with their distance from the transmitter (Fig. 1a). This method reduces the effect of the superficial geophysical noise on larger receiver dipoles, it uses three investigation depths without switching and the size of the arrays remains limited to a value equivalent to that of the investigation depth (Panissod et al., 1997; Dabas, 2009). The collected apparent resistivity data are presented in the form of 3 maps corresponding to the different measuring dipoles and the recorded anomalies are interpreted in terms of possible buried archaeological structures. These maps represent the cumulative volume contribution of the soil from top to the three increasing depth of investigations.

In special cases where a more quantitative interpretation is necessary, the 3D resistivity inversion of the data can provide additional information about the burial depth and the depth extent of buried relics (Papadopoulos et al., 2007). The resistivity inversion scheme starts by lineariz-
ing the inverse problem and minimizing the data misfit in the least squares sense. The instability and non-uniqueness of the inverse problem is approached by imposing specific smoothness constraints (Sasaki, 1994). In this work, a relatively small region from a wider area previously explored by the ARP method was selected and re-processed by a 3D resistivity inversion algorithm in order to acquire a more quantitative model of the buried archaeological structure.

A large geophysical exploration was undertaken at Andilly archaeological site in France. An area of more than 17,500 square meters was investigated with the ARP method. The transmitting and receiving electrode configuration for the mobile device is shown in figure 1(a). The data were acquired along 1 m parallel spaced profiles at 0.2 m sampling rate. A Kriging interpolation algorithm was used to grid the data at 0.5 m spacing in both horizontal directions. Two large rectangular high apparent-resistivity anomalies are outlined at the center and at the east of the surveyed area. Some rectangular structures are also visible at the southeastern corner of the region. These anomalies are registered in all the maps that correspond to the different length of the receiving dipoles.

A 10 × 10 m area, which encloses the connection of two vertical structures (Fig. 1b), was extracted from the re-sampled original data. These data were used as input into the 3D resistivity inversion algorithm. The geometry and the coordinates of the electrodes were appropriately configured by assigning to every different measurement a unique four-electrode combination (Fig. 2a), based on the electrode configuration of figure 1(a) and the step interval of 0.5 m along the X and Y axes. This data transformation resulted in 1873 different electrode positions which cover the 10 × 10 m grid and 1323 measurements that described the 3D apparent resistivity variation.

The subsurface resistivity distribution was described by 3750 (25 × 26 × 5) cubic parameters. The horizontal dimensions and the thickness of these parameters were set equal to the 0.5 m re-sampled step interval of the ARP survey. The thick lines in figure 2 (b, c) indicate the parameters of constant resistivity into which the area was divided. On the other hand, the thin lines show the distribution of 3D elements that form a specific parameter. These elements formulate the 3D Finite Element Mesh (FEM) which was used to calculate the synthetic apparent resistivity data in the inversion procedure. A 3D Finite Element inversion algorithm was used to process and invert the ARP data. The inversion procedure is based on a smoothness constrained Gauss-
Newton algorithm. The Active Constraint Balancing (ACB) method was also incorporated in the inversion in order to enhance the least-squares resolving power and stability (Yi et al., 2001). The values of 0.01 and 1 were used for the lower (\(\lambda_{\min}\)) and the upper (\(\lambda_{\max}\)) bounds of the Lagrange multiplier respectively. These values are necessary for the spread function analysis and they were defined a priori by conducting some initial test inversions in the data set.

The inversion algorithm converged to a final 3D resistivity model after 7 iterations and \(RMS = 2.3\%\). Horizontal depth slices every 0.25 m were extracted from the inversion model and they are presented in figure 3. The superficial layer \((Z = 0.375\ m)\) has large resistivity variations. The two vertical structures appear from the depth \(Z = 0.625\ m\) and reach the depth of \(Z = 1.625\ m\) below ground surface. The maximum burial depth of the structures does not exceed \(2.3\ m\). Some faint remnants below that depth are due to the limited resolution of the surface electrical survey and the smoothness constraints imposed in order to stabilize the inversion procedure.

The 3D inversion managed to reconstruct the subsurface resistivity and identify the two vertical structures. Additional information concerning the burial depth and the depth extent of the archaeological features was extracted by the inversion. Compared to the three ARP apparent resistivity maps, 3D inversion can provide a more quantitative insight of the data. This information is very important to the archaeological team in order to plan more accurately future excavation projects.

On the other hand, the application of the ARP method involves the collection of a huge amount of data. This actually prohibits the 3D inversion of such data sets. So the inversion can only be applied in relatively small areas that exhibit a special interest and for which extra information concerning more quantitative characteristics needs to be determined.
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References


